

1 **Natural archives of long-range transported contamination at the remote lake Letšeng-la Letsie,**
2 **Maloti Mountains, Lesotho**

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21

22 **Abstract**

23 Naturally accumulating archives, such as lake sediments and wetland peats, in remote areas may be
24 used to identify the scale and rates of atmospherically deposited pollutant inputs to natural
25 ecosystems. Co-located lake sediment and wetland cores were collected from Letšeng-la Letsie, a
26 remote lake in the Maloti Mountains of southern Lesotho. The cores were radiometrically dated and

27 analysed for a suite of contaminants including trace metals and metalloids (Hg, Pb, Cu, Ni, Zn, As), fly-
28 ash particles, stable nitrogen isotopes, polycyclic aromatic hydrocarbons (PAHs) and persistent organic
29 pollutants such as polychlorinated biphenyls (PCBs), polybrominated flame retardants (PBDEs) and
30 hexachlorobenzene (HCB). While most trace metals showed no recent enrichment, mercury, fly-ash
31 particles, high molecular weight PAHs and total PCBs showed low but increasing levels of
32 contamination since c.1970, likely the result of long-range transport from coal combustion and other
33 industrial sources in the Highveld region of South Africa. However, back-trajectory analysis revealed
34 that atmospheric transport from this region to southern Lesotho is infrequent and the scale of
35 contamination is low. To our knowledge, these data represent the first palaeolimnological records and
36 the first trace contaminant data for Lesotho, and one of the first multi-pollutant historical records for
37 southern Africa. They therefore provide a baseline for future regional assessments in the context of
38 continued coal combustion in South Africa through to the mid-21st century.

39

40 **Keywords:** contamination; fly-ash; lake sediments; Lesotho; long-range transport; persistent organic
41 pollutants; trace metals; wetlands

42

43 **1. Introduction**

44

45 Southern Africa possesses a wealth of natural landscapes from the mountainous terrain and upland
46 tarns of the Drakensberg and Great Escarpment to coastal lagoons, and from open savannah to
47 evergreen and cycad forests. These wide-ranging habitats have resulted in a diverse endemic fauna
48 and flora and globally important biodiversity hotspots (Myers et al., 2000; CEPF, 2011). However, a
49 range of pressures detrimentally affect these natural landscapes, and especially freshwater
50 ecosystems. These include expanding urbanisation; an increase in water abstraction and invasive
51 species (Driver et al., 2012); a rapid increase in the release of pollutants to the environment from
52 industrial emissions (Monna et al., 2006; Josipovic et al., 2011) including a continued reliance on fossil-

53 fuels (Marais et al., 2019); ongoing use of DDT to control malarial vectors (Sereda and Meinhardt,
54 2005; Humphries, 2013); direct inputs to waters from acid mine drainage (Tutu et al., 2008; McCarthy,
55 2011) and agricultural practices. Freshwaters in southern Africa are therefore experiencing a period
56 of unprecedented threat. This is likely to be exacerbated by predicted changes to precipitation and
57 increases in temperature of 0.3-3.6 °C by 2036-2065 over a 1961-1990 baseline (Davis, 2011; SAWS,
58 2017), making southern Africa one of the world's most climatically vulnerable regions. There are,
59 therefore, significant implications for aquatic biodiversity, human health, and ecosystem service
60 provision to dependent rural communities.

61

62 Lesotho is a land-locked country, entirely surrounded by South Africa and over 80% of the country lies
63 over 1800 m above sea level (a.s.l) (Schmitz and Rooyani, 1987). The Maloti Mountains, part of the
64 Drakensberg range and considered part of the Afromontane zone, form a high alpine basalt plateau
65 across the southern and eastern parts of the country, incised by deep valleys which drain into the
66 Senqu River, known as the Orange River in South Africa (Schmitz and Rooyani, 1987; Marake et al.,
67 1998). The population of Lesotho is estimated at just over 2 million (Lesotho Bureau of Statistics, 2016)
68 and Maseru, the capital, along with the main agricultural areas lie in the west (Letšela, 2008). The
69 economy is based on agriculture, livestock, manufacturing and mining, while water and diamonds are
70 the significant natural resources. The Lesotho Highlands Water Project (LHWP), which started in 1986,
71 is designed to store and transfer water from the Senqu/Orange River system to South Africa's Free
72 State province and Johannesburg (Letšela, 2008). Water from Lesotho provides around 35% of the
73 water supplied by the Vaal Water Supply System, which supplies the Gauteng and Mpumalanga urban
74 and industrial complex, including mines, industries and most coal-fired power stations (Le Maitre et
75 al., 2018). Completion of the first phase of the LHWP has contributed significantly towards Lesotho's
76 self-sufficiency in the production of electricity. However, electricity demand currently outstrips
77 supply, and Lesotho imports 37% of its annual electricity needs from South Africa (Thamae et al.,
78 2015). Peak demand is expected to grow to over 500 MW by 2050 (Senatla et al., 2018).

79

80 By contrast, South Africa is one of the major industrial countries of the southern hemisphere and the
81 largest industrialised economy in Africa (Venter et al., 2017). It has significant mining and metallurgical
82 activities and is one of the top coal-producing countries in the world (Chelgani, 2019). It has rich coal
83 deposits in the north-east of the country and as a consequence the majority of South Africa's coal-
84 fired plants are located here, in the Highveld area of Mpumalanga province. One estimate suggests
85 South Africa accounted for 43% of the total generating capacity for the continent in 2012 (Marais et
86 al., 2019). Over 85% of all electricity generated in South Africa in 2016 was from coal (StatsSA, 2018)
87 and there is a programme of large, coal-fired power station construction in the Highveld coalfields and
88 in the Lephalale area of Limpopo province, near the border with Botswana. The scarcity of major
89 combustion sources in other parts of the country results in strong spatial gradients of emissions from
90 east to west and from north to south (Josipovic et al., 2011). However, deposition monitoring for fossil-
91 fuel associated contaminants is limited and so long-range transported pollution gradients are assumed
92 to follow patterns in sulphur and nitrogen deposition modelling studies (Lamarque et al., 2013), which
93 indicate highest deposition in the north-east, declining to the south and west. By contrast, local
94 sources of pollutants, such as agricultural pesticide use, and biomass burning due to prescribed
95 management or accidental fires, may affect catchments anywhere in the country.

96

97 Contamination of South African aquatic environments by trace metals and persistent organic
98 pollutants (POPs) is widely acknowledged. The presence of metals and organochlorine pesticides
99 (principally DDT and its derivatives, but also polychlorinated biphenyls (PCBs), dioxins and furans
100 (PCDD/Fs)) have been reported for riverine plants, fish, sediments and otters (Mason and Rowe-Rowe,
101 1992; Roux et al., 1994; Grobler et al., 1996; Vosloo and Bouwman, 2005) with “elevated”
102 accumulations of metals in fish from the Olifants River, Mpumalanga (Kotze et al., 1999). In the
103 lacustrine environment, metals and organohalogen compounds (including DDT, PCBs and brominated
104 flame retardants such as polybrominated diphenyl ethers (PBDEs)) have been reported in fish,

105 crayfish, sediments and aquatic birds (Greichus et al., 1977; Bezuidenhout et al., 1990; Sereda and
106 Meinhardt, 2005; Bouwman et al., 2008; Nakayama et al., 2010; Wepener et al., 2012; Humphries,
107 2013). Evidence for bioconcentration in higher trophic levels resulting in eggshell thinning has been
108 shown for the aquatic-feeding African darter (*Anhinga rufa*) due to intake of DDE (a
109 dehydrochlorinated product from DDT) and PCBs (Bouwman et al., 2008). Yet, despite this evidence
110 for adverse effects, to our knowledge, there are no equivalent data for Lesotho and very few data
111 exist on the scale and rate of change of contaminant inputs to freshwater systems anywhere in
112 southern Africa.

113

114 In other parts of the world, naturally accumulating archives such as lake sediments and wetland peats
115 have been used extensively to identify the impacts of environmental pollutant inputs to natural
116 ecosystems. These archives have been used to identify directions of change (e.g. improvement or
117 deterioration), the scale of anthropogenic enrichment of contaminants over natural baselines and,
118 perhaps most importantly, the changing rate of inputs. However, while high-resolution lake sediment
119 records of trace metals and POPs are widely used, particularly in Europe and North America, in
120 southern Africa such techniques have so far only been applied to a very limited number of sites
121 although this is, in part, due to the rarity of natural, permanent standing waters. To date, only two
122 eutrophic urban lakes, Zeekoevlei in the western Cape (Das et al., 2008a,b) and North End Lake in Port
123 Elizabeth (García-Rodríguez et al., 2007) and a single peat core from the Witwatersrand industrial area
124 (McCarthy and Venter, 2006) have been analysed in this way. Hence, there are no assessments of
125 regional contaminant patterns and no record of atmospherically deposited contamination in remote
126 and non-urban freshwater systems.

127

128 Remote mountain lakes can act as sentinels for environmental change as their contaminant sources
129 are usually solely atmospherically derived. Their archives can therefore provide a means to determine
130 regional patterns and baselines for future change not usually possible in lowland systems. In this study

131 we assess the historical record of contamination in the remote lake, Letšeng-la Letsie, located in the
132 Maloti Mountains of Lesotho. This is the only RAMSAR Convention site (www.ramsar.org) in Lesotho,
133 designated as a high altitude wetland, and for IUCN-classified vulnerable bird species including
134 Wattled cranes (*Grus carunculatus*), Lesser kestrel (*Falco naumanni*), Bald ibis (*Geronticus calvus*), Cape
135 vulture (*Gyps coprotheres*), and the Yellow-breasted pipit (*Anthus chloris*). The Maloti minnow
136 (*Pseudobarbus quathlambae*), which may occur at the site, is considered critically endangered and
137 rare (Motanya, 2003). While potential threats from over-grazing and catchment soil erosion are
138 recognised (Matope, 2014) there is no information on the scale of contamination.

139

140 Letšeng-la Letsie provides a unique opportunity to study a paired lake sediment and wetland archive
141 in Lesotho. The aims of this study were to analyse these historical records for a suite of contaminants,
142 including trace metals and metalloids (Hg, Pb, Cu, Ni, Zn, As) (hereafter ‘trace metals’), spheroidal
143 carbonaceous fly-ash particles (SCPs), stable nitrogen isotopes and organic pollutants including PAHs
144 and POPs, to determine the scale of contamination at this remote site, and its likely sources. In so
145 doing, the study provides the first palaeolimnological records for Lesotho and, more broadly, the first
146 high resolution, multi-pollutant records for southern Africa. An additional aim, therefore, is to add to
147 the sparse southern hemisphere dataset for these pollutants and provide a baseline for future regional
148 assessments in the context of continued coal combustion in South Africa through to the mid-21st
149 century. Given the contrast in the scale of emissions between South Africa and Lesotho, we
150 hypothesise that any contaminants in this remote area are likely derived by the long-range transport
151 of emissions from industrial sources in north-eastern South Africa.

152

153

154 **2. Methods**

155

156 *2.1. Site description*

157

158 Letšeng-la Letsie (30°18'39.7"S; 28°09'59.3"E) is a 0.38 km² lake in the Maloti Mountains of Lesotho,
159 lying at 2400 m a.s.l., about 200 km south-east of Maseru and close to the border between Lesotho
160 and Eastern Cape Province, South Africa (Fig. 1). The 413.9 km² catchment area was declared the
161 Letšeng-la Letsie Protected Area by the Principal Chief of the Area together with the Prime Minister
162 of Lesotho in 2001 and formed part of the 'Conserving Mountain Biodiversity in Southern Lesotho'
163 area (CMBSL) (Motanya, 2003). The lake has a mean depth of about 1 m and a maximum recorded
164 depth of 1.4 m (Fig. 2). It was impounded in 1968 by a small dam across the Mohlakeng River, which
165 forms the lake's outflow. This flows via the Quthing River to form a tributary of the Senqu River
166 becoming the Orange River in South Africa. The dam was already reported to be in a poor state of
167 repair in 1988 (Skofteland, 1988) and although the water-level of the lake undoubtedly remains raised,
168 the effectiveness of the dam was limited during the field sampling in 2018 with a steady flow of water
169 by-passing it on either side.

170

171 The catchment is bounded by mountains that rise to 2820 m a.s.l giving a relief of 420 m. The dominant
172 geology is basaltic with dolerite intrusions, while the valley floor including the lake have recent alluvial
173 deposits. The region falls within the Drakensberg Alpine Complex (Carbutt and Edwards, 2004; 2006),
174 with Afroalpine and Afroalpine vegetation dominated by grasses, sedges and woody shrubs. There
175 are no trees, as the region is situated above the tree-line (Nüsser and Grab, 2002). Letšeng-la Letsie is
176 located within an area of the Maloti-Drakensberg Mountains known as a "biodiversity hotspot" with
177 over 30% endemism (Lesotho Government National Environment Secretariat, 2000). These include
178 the mountain pipit (*Anthus hoeschi*), while another six species including the Bush blackcap (*Lioptilus*
179 *nigricapillus*), Buff-streaked chat (*Saxicola bifasciata*), Rudd's lark (*Heteromirafra ruddi*), Orange
180 breasted rockjumper (*Chaetops aurantius*) and Drakensberg siskin (*Serinus symonsi*) are near
181 endemic. The only endemic mammal and amphibian at the site are the ice rat (*Otomys slogetti*) and
182 the aquatic river frog (*Rana vertebralis*), respectively.

183

184 Winters can be extreme with temperatures commonly -5 °C and a recorded minimum of -22 °C. Snow
185 falls are common in the austral winter months although snow, sleet or hail can occur at any time of
186 the year (Grab and Nash, 2010; Grab and Linde, 2014). Maximum temperatures of 26 °C have been
187 recorded in January (Lesotho Ministry of Natural Resources, 1996). Highest rainfall occurs at the
188 mountain summits which receive over 1000 mm per annum (Sene et al., 1998). More than 80% of
189 rainfall is received between October and March, and on average, the area receives about 800 mm
190 rainfall each year (Sene et al., 1998; Borg, 2012). Lesotho is amongst the windiest countries in Africa
191 (Mpholo et al., 2012). The strongest winds are experienced in late-winter due to the frontal systems
192 associated with the westerlies and lowest in the wet summer months (Grab, 2002a; 2010). Little
193 climate data exist for this area of Lesotho, but in the Sani Valley to the east, summer winds dominate
194 from the east and northeast, bearing moisture from the Indian Ocean, while the winter winds are
195 predominantly north-westerly (Sene et al., 1998; Grab, 2010).

196

197 *2.2. Sediment and wetland cores*

198

199 Three parallel lake sediment cores (LETS1-3) were collected from the deepest part of the lake (1.4 m)
200 to the east of the outflow area (Fig. 2) from an inflatable boat. The 26 – 30 cm cores were taken using
201 an HTH gravity corer fitted with a perspex tube of internal diameter 86 mm (Renberg and Hansson,
202 2008). The cores were extruded vertically in the field at 0.5 cm intervals. All analyses except those for
203 organic compounds were undertaken on LETS2. The core tube and extruding equipment for LETS3,
204 used for organic compound analyses were *n*-hexane-washed and wrapped in *n*-hexane-cleaned
205 aluminium foil prior to use. The cores were kept cold and dark, and were frozen (-20 °C) as soon as
206 possible. Replicate wetland sediment monoliths were also collected from two locations to the north-
207 west of the lake. The monoliths, ranging from 23 - 34 cm length, were collected from exposed vertical
208 sections in drainage channels using strong plastic tubing following standard techniques (e.g. De

209 Vleeschouwer et al., 2010). Before sampling, the sediment face was cut back to expose
210 uncontaminated and un-oxidised material. Once collected, the monolith was wrapped and stored at
211 4 °C before sub-sampling at 0.5 cm intervals. Monolith LETS-W3 was used for all contaminant analyses.

212

213 2.3. Lithostratigraphic analyses and core chronologies

214

215 Lithostratigraphic analyses (bulk density; water content; loss-on-ignition at 550°C (LOI₅₅₀) and 950 °C
216 (LOI₉₅₀) as an estimate of organic matter content and carbonate content respectively) were
217 undertaken on each sample following standard methods (Dean, 1974; Heiri et al., 2001). Core samples
218 were analysed for ²¹⁰Pb, ²²⁶Ra, ¹³⁷Cs and ²⁴¹Am by direct gamma assay using ORTEC HPGe GWL series
219 well-type coaxial low background intrinsic germanium detectors. Lead-210 was determined via its
220 gamma emissions at 46.5 keV, and ²²⁶Ra by the 295 keV and 352 keV gamma rays emitted by its
221 daughter isotope ²¹⁴Pb following three weeks storage in sealed containers to allow radioactive
222 equilibration for ²²²Rn and ²²⁶Ra. Cesium-137 and ²⁴¹Am were measured by their emissions at 662 keV
223 and 59.5 keV, respectively. The absolute efficiencies of the detector were determined using calibrated
224 sources and sediment samples of known activity. Corrections were made for self-absorption of low
225 energy gamma rays within the sample (Appleby et al., 1992). Unsupported ²¹⁰Pb, from atmospheric
226 deposition, was calculated by subtracting supported ²¹⁰Pb (which derives from *in situ* decay of ²²⁶Ra)
227 from total ²¹⁰Pb. Final sediment chronologies were determined from ²¹⁰Pb records using constant rate
228 of supply (CRS) or constant initial concentration (CIC) models (Appleby, 2001), in combination with
229 ¹³⁷Cs and ²⁴¹Am profiles. Derived and interpolated sediment accumulation rates (g cm⁻² yr⁻¹) were used
230 to convert contaminant concentrations into fluxes.

231

232 2.4. Sediment and wetland core geochemistry and trace metal analysis

233

234 Weighed (4 d.p.) freeze-dried milled sediment (~2 g) was measured using an X-ray fluorescence
235 spectrophotometer (Rigaku NEX CG EDXRF) for trace metals (Pb, Cu, Ni, Zn, As) and other geochemical
236 elements. A certified reference sediment sample of similar mass was included in each analytical run
237 (Canadian Certified Reference Materials Project; LKSD-2) with mean recovery rates of 93.6 (Cu) to
238 102.5% (Zn). For mercury (Hg) analysis, 0.2 g of freeze-dried samples were weighed into a 50 mL
239 polypropylene DigiTUBE (SCP Science). 8 mL of 'aqua regia' (nitric and hydrochloric acids in a molar
240 ratio of 1:3) were added to each and gradually heated on a hotplate to 100 °C to avoid violent reaction.
241 The sample was then digested for another 1.5 h and allowed to cool. The digested solution was diluted
242 to 50 mL using distilled deionised water. Standard reference stream sediment (GBW07305; certified
243 Hg value $100 \pm 10 \text{ ng g}^{-1}$; our measured mean value 100.3 ng g^{-1} ; RSD = 4.5 ng g^{-1} ; N=15) and sample
244 blanks were digested with every 20 samples. Mercury concentrations were measured by cold vapour-
245 atomic fluorescence spectrometry (CV-AFS) following reduction with SnCl_2 . Standard solutions and
246 quality control blanks were measured after every five samples to monitor measurement stability.

247

248 Trace element enrichment factors (EFs) provide a means by which to determine elevation above
249 uncontaminated background levels and hence estimate anthropogenic inputs to sediments (e.g. Chen
250 et al., 2007). EFs were calculated by normalising trace element concentrations to those of a
251 conservative geochemical element within the measured sample and then to the same ratio within
252 basal (uncontaminated) sediments using the following equation:

253

$$254 \quad \text{EF} = \left(\frac{[M_{\text{samp}}]}{[R_{\text{samp}}]} \right) / \left(\frac{[M_{\text{back}}]}{[R_{\text{back}}]} \right) \quad (1)$$

255

256 where M_{samp} is the metal concentration in the sample for which the EF is being calculated, R_{samp} is the
257 concentration of the reference element (here, Ti), while M_{back} and R_{back} are the equivalent
258 concentrations in the basal sample. EFs of 1 - 3 represent minor anthropogenic contamination; 3 - 5

259 moderate; 5 - 10 moderate to severe; 10 - 25 severe; 25 - 50 very severe and > 50 extremely severe
260 contamination (Chen et al., 2007).

261

262 *2.5. Spheroidal carbonaceous fly-ash particles (SCPs)*

263

264 Spheroidal carbonaceous particles (SCPs) are a component of fly-ash, produced solely from the high-
265 temperature combustion of coal-series and fuel oils (Rose, 2001). They have no natural sources and
266 therefore provide an unambiguous indicator of atmospherically deposited industrial contamination.
267 Although particulate, SCPs are able to travel hundreds to thousands of kilometres (Rose et al., 2012)
268 with diffusing industrial plumes before being dry or wet deposited depending on meteorological
269 conditions. For both lake and wetland core samples, SCP analysis involved sequential treatments of
270 nitric, hydrofluoric and hydrochloric acids to remove organic, siliceous and carbonate fractions,
271 resulting in a suspension in water (Rose, 1994). A known fraction of this suspension was then
272 evaporated onto a coverslip, mounted onto a glass slide, and the number of SCPs counted using a light
273 microscope at 400 times magnification. Standard criteria for SCP identification were followed (Rose,
274 2008). SCP concentrations were calculated as the number of particles per gram dry mass of sediment
275 (gDM^{-1}) and SCP fluxes calculated as the product of SCP concentration and bulk dry sediment
276 accumulation rate (number of particles per cm^2 per year; $\text{cm}^{-2} \text{yr}^{-1}$). Analytical blanks and SCP reference
277 material (Rose, 2008) were included with all sample digestions. The detection limit for the technique
278 is typically less than 100 gDM^{-1} and calculated concentrations generally have an accuracy of c. ± 45
279 gDM^{-1} .

280

281 *2.6. Stable isotopes of nitrogen*

282

283 Samples were analysed for their $\delta^{15}\text{N}$ signatures using a Flash EA 1112 and a Delta V continuous-flow
284 gas-isotope-ratio mass spectrometer (Thermo Scientific®) in the Bloomsbury Environmental Isotope

285 Facility, UCL. Samples of 10-20 mg were accurately weighed into tin capsules which were dropped into
286 a furnace held at 1020 °C and burned with a pulse of oxygen at about 1800 °C. The released gases
287 were carried in a stream of helium through an oxidant (granular chromium oxide combustion reagent)
288 to complete combustion; silvered cobalt chemical traps to remove sulphur and halogens; a copper
289 reduction furnace to trap excess oxygen and to convert the formed NO_x to N₂ for analysis; a
290 magnesium perchlorate drying tube which also contained some Carbosorb to remove CO₂. Finally, the
291 gases pass through a gas chromatograph (GC) column before passing through a Thermal Conductivity
292 Detector (TCD, to obtain TN%) to then enter the mass spectrometer via a CONFLO IV interface. The
293 results were calibrated against laboratory and international standards (i.e. IAEA-600, IAEA-N1, IAEA-
294 N2, USGS40, OEA alanine). Nitrogen isotope ratios are reported in the conventional delta-notation, in
295 per mil (‰) with respect to atmospheric N₂ (Air).

296

297 *2.7. Organic pollutants*

298

299 *2.7.1. Sample extraction and fractioning of PAHs and POPs*

300 The analytical procedure for the determination of polychlorinated biphenyl (PCBs), brominated
301 diphenyl ethers (BDEs), hexachlorobenzene (HCB) and polycyclic aromatic hydrocarbons (PAHs) was
302 based on a UNEP method (1992) as modified by Bicego et al. (2006). Approximately 15 g dried
303 sediment was Soxhlet extracted for 8 h with 80 mL of a *n*-hexane and dichloromethane (DCM) mixture
304 (1:1, v/v), copper (for the removal of elemental sulphur), boiling glass spheres and 100 µL of a
305 surrogate standards mixture containing acenaphthene-d₁₀, phenanthrene-d₁₀ and chrysene-d₁₂ (2.5
306 ng µL⁻¹) for PAHs, and PCB 103 and PCB 198 (1 ng µL⁻¹) for POPs. The extract was reduced to 4 mL using
307 rotary evaporation and divided into two portions of 2 mL. The first portion was subjected to a clean-
308 up and fractionation procedure using a glass column containing 3.2 g of silica, 1.8 g of alumina (both
309 5% deactivated with Milli-Q® water) and sodium sulphate. The column was first eluted with 10 mL of
310 *n*-hexane to remove a first fraction containing aliphatic hydrocarbons (data not presented), and then

311 with 15 mL of a DCM and *n*-hexane mixture (3:7, v/v), to remove a second fraction (PAHs).
312 Subsequently, these fractions were concentrated by rotary evaporation and transferred to calibrated
313 glass vials with a final volume of 250 μL . Before instrumental analysis, 50 μL of the internal standard
314 benzo[*b*]fluoranthene- d_{12} (2.5 $\text{ng } \mu\text{L}^{-1}$) was added to each PAH extract and adjusted to 250 μL . The
315 second 2 mL portion was purified using a glass column containing 3.2 g of 5% deactivated aluminium
316 oxide, and subsequently eluted with 20 mL of a DCM and *n*-hexane mixture (3:7, v:v), to obtain the
317 POPs fraction. The eluate was spiked with 50 μL of the internal standard tetra-chloro-*m*-xylene (TCMX,
318 1 $\text{ng } \mu\text{L}^{-1}$) and concentrated to 250 μL for POPs analysis.

319

320 *2.7.2. Instrumental analysis of PAHs and POPs*

321 The instrumental analyses for PCBs, PBDEs, HCB and PAHs were performed by Agilent 7890A gas
322 chromatography (GC) coupled with an Agilent 5975C inert MSD with a Triple-Axis Detector Mass
323 Spectrometer (MS). The GC column was an HP-5 fused silica column (length: 30 m, 0.25 mm ID, film
324 thickness: 0.25 μm). The oven temperature for PCBs, PBDEs and HCB was programmed to begin at 75
325 $^{\circ}\text{C}$ for 3 min, increasing at 15 $^{\circ}\text{C } \text{min}^{-1}$ up to 150 $^{\circ}\text{C}$, then at 2 $^{\circ}\text{C } \text{min}^{-1}$ up to 260 $^{\circ}\text{C}$, and finally at 20 $^{\circ}\text{C}$
326 min^{-1} to up to 300 $^{\circ}\text{C}$ (holding for 10 min), in a total time of 75 min. The oven temperature for PAHs
327 was: 40 $^{\circ}\text{C}$ to 60 $^{\circ}\text{C}$ at 20 $^{\circ}\text{C } \text{min}^{-1}$, then to 250 $^{\circ}\text{C}$ at 5 $^{\circ}\text{C } \text{min}^{-1}$ and, finally, to 300 $^{\circ}\text{C}$ at 6 $^{\circ}\text{C } \text{min}^{-1}$ (held
328 for 20 min). For all analyses, the injector temperature was adjusted to 280 $^{\circ}\text{C}$, splitless injection mode
329 was adopted and the detector and ion source temperatures were adjusted to 300 $^{\circ}\text{C}$ and 230 $^{\circ}\text{C}$,
330 respectively. Helium was used as the carrier gas at a constant flow rate of 1 $\text{mL } \text{min}^{-1}$ while data
331 acquisition was obtained using Selected Ion Monitoring (SIM) mode, and quantification was based on
332 the integration of specific fragment ion (m/z) peaks area using an Agilent Enhanced Chemstation
333 (G1701 CA) program. The PAHs were identified by matching the retention time and specific m/z with
334 the results obtained from the standard mixtures (Z-014G-FL, AccuStandard, USA), and with a
335 calibration curve ranging from 0.10 to 2.00 $\text{ng } \mu\text{L}^{-1}$ ($r^2 > 0.995$). The concentration of individual PCBs
336 and PBDEs was obtained through the internal standard peak area method and an 8-point analytical

337 curve for individual components (1, 5, 10, 20, 50, 80 and 100 $\mu\text{g L}^{-1}$, $r^2 > 0.995$) from the standard
338 mixtures (PCBs: C-WNN and C-WCFS; PBDEs: BDE-CM, from AccuStandard, USA). The final
339 concentrations were expressed as ng g^{-1} in dry sediment for all compounds detected.

340

341 The following compounds were investigated: (i) low-chlorinated PCBs: 8, 18, 28, 31, 33, 44, 49, 52,
342 56/60, 66, 70, and 74; (ii) high-chlorinated PCBs: 87, 95, 97, 99, 101, 105, 110, 114, 118, 123, 128, 132,
343 138, 141, 149, 151, 153, 156, 157, 158, 167, 170, 174, 177, 180, 183, 187, 189, 194, 195, 201, 203, 206,
344 and 209; (iii) BDEs: 28, 47, 99, 100, 153, 154 and 183; (iv) hexachlorobenzene (HCB); (v) PAHs: (a) alkyl-
345 PAHs including methyl (C_1), dimethyl (C_2) and trimethyl (C_3) naphthalenes, methyl phenanthrenes; (b)
346 unsubstituted PAHs (2–3 rings) such as naphthalene (C_0), phenanthrene (C_0), anthracene, fluorene,
347 acenaphthene and acenaphthylene; (c) PAHs (4–6 rings) such as fluoranthene, pyrene,
348 benz[*a*]anthracene, chrysene, benzo[*b+k+j*]fluoranthene, benzo[*a+e*]pyrene, perylene, indeno[1,2,3-
349 *c,d*]pyrene, dibenz[*a,h*]anthracene and benzo[*g,h,i*]perylene.

350

351 *2.7.3. Quality assurance for organic pollutants*

352 Quality assurance was based on blank extraction (in 15 g of sodium sulphate) and the recoveries of
353 surrogate standards. Procedural blanks were undertaken for each group of ten samples and when
354 necessary (e.g. values greater than 3 times detection limits), the values found in the blanks were
355 removed from the samples. The mean PAH surrogate recoveries in the extracted samples were: (i)
356 PAH: 55 ± 10 % for acenaphthene- d_{10} , 78 ± 14 % for phenanthrene- d_{10} and 83 ± 12 % for chrysene- d_{12} ,
357 (ii) PCB: 95 ± 22 % for PCB 103 and 73 ± 16 % for PCB 198. Surrogate recoveries in the samples were
358 in the acceptable range (40 - 120%; Denoux et al., 1998) for 100% of samples analysed. Precision
359 ranged from 3.2 to 10.4% for individual PAHs and from 0.2 to 11.4 % for individual PCBs. Method
360 accuracy was checked by the extraction of duplicate standard reference material (SRM) for sediment
361 IAEA 417 (International Atomic Energy Agency, Vienna, 1999). The results were within the upper and
362 lower ± 30 % confidence interval reference values for 81% of PAHs and 91 % of PCBs in accordance

363 with criteria proposed by Wade and Cantillo (1994). The instrumental detection limits named as
364 detection limits (DL) adopted in this study were 0.50 ng g^{-1} for the PAHs and 0.03 ng g^{-1} for the PCBs
365 and PBDEs. These data are based on the lowest sensitive PAH / PCB concentration ($0.03 \text{ ng } \mu\text{L}^{-1} / 0.002$
366 $\text{ng } \mu\text{L}^{-1}$) multiplied by the final extracted volume ($250 \text{ } \mu\text{L}$) and divided by the sediment weight (15 g)
367 before extraction. The analysis of duplicate samples was used to verify precision which was $< 15 \%$ for
368 the individual compounds analysed.

369

370 *2.8. Back trajectory analysis*

371

372 The origin of air flowing over Letšeng-la Letsie was investigated with NOAA's Hybrid Single-Particle
373 Lagrangian Integrated Trajectory (HYSPPLIT) model (Stein et al., 2015), using global NCEP reanalysis
374 data. Back trajectories were run for 72 hours, with point of origin 1500 metres above ground level at
375 Letšeng-la Letsie. Trajectories were calculated twice a day (at 00:00 and 12:00 UTC) at ten-year
376 intervals, from 1950 to 2010. Predominant transport pathways were identified using the HYSPPLIT
377 model's trajectory cluster analysis, which minimises the total spatial variance between the mean of a
378 cluster and all the trajectories assigned to that cluster (Draxler et al., 2012).

379

380 *2.9. Emissions from coal combustion*

381

382 The amount of coal burnt per annum for electricity generation in South Africa was obtained from
383 Eskom's annual reports, to determine to what extent the trends in pollutants observed in the lake
384 sediments can be ascribed to trends in emissions. Eskom's stations accounted for about 75% of all coal
385 used for power generation in South Africa in the 1940s. In the late-1940s, Eskom procured several
386 additional power stations, so that by the 1950s the bulk of coal-fired generation capacity in South
387 Africa was operated by them. The coal used for power generation accounts for the majority of coal
388 use in South Africa and, between 2006 and 2015, coal-fired power stations consumed on average 66%

389 of all coal consumed in the country (Ratshomo and Nembaha, n.d.). Sulphur dioxide (SO₂) emissions
390 were obtained from Eskom's annual reports (Eskom, 2019); SO₂ emissions were calculated with a
391 station-specific emission factor, based on the sulphur content of the coal, from 1927 to 1981, and with
392 mass balance calculations using daily measurements of the sulphur content in the coal thereafter.
393 Mercury emissions were calculated from mass balance, based on the mercury in coal measurements
394 reported in Garnham and Langerman (2016) for currently operating stations and a national mercury-
395 in-coal concentration average value of 0.31 ppm (UN Environment, 2017) for older stations for which
396 there are no measurements. It was assumed that the abatement technology installed at a power
397 station reduces mercury emissions by 25% for an electrostatic precipitator or cyclone, and by 50% for
398 a fabric filter plant (UN Environment, 2017).

399

400 **3. Results**

401

402 *3.1. Core chronologies*

403

404 The radiometric dating for both the lake and wetland cores are shown in Supplementary Information
405 (SI1). The ²¹⁰Pb and ¹³⁷Cs records provide chronologies extending to 1933 ± 16 (85 years) and 1899 ±
406 23 (119 years) respectively. Both show records of continuous accumulation before and after the
407 construction of the impoundment in 1968. The lake sediment core (LETS2) shows little change in
408 sediment accumulation as a result of impoundment although the rate appears to become more stable
409 afterwards. For the wetland core (LETS-W3) there is a continuous trend of increasing sediment
410 accumulation rate from the base of the core through to the present, which appears to accelerate from
411 the late-1980s. The dry weight and loss-on-ignition (550 °C and 950 °C) profiles for the sediment cores
412 LETS2 and LETS3 show excellent agreement (Supp. Info. SI1) and were used to apply dates to the
413 otherwise undated LETS3 core.

414

415 3.2. Trace metals, SCPs and stable isotopes of nitrogen

416

417 The concentrations of the trace metals and metalloids (Ni, Cu, Zn, Pb, Hg and As; hereafter 'trace
418 metals') are shown in Fig. 3. There is little change in metal concentration over the period covered by
419 the cores. Only Hg shows evidence for an increase with concentrations in the lake core increasing
420 unidirectionally from mid-30s ng g^{-1} at the base of core to $> 50 \text{ ng g}^{-1}$ in recent sediments. Due to this
421 lack of significant change in metal concentration, metal fluxes follow the historical trends in sediment
422 accumulation rates (Supp. Info. SI2). As a result, metal fluxes for LETS2 show some degree of variability
423 through the 1950s and 1960s but become more stable from the start of the 1970s at the time of the
424 impoundment. By contrast, the increasing sedimentation rates in the wetland core result in increasing
425 metal fluxes from the start of the record, such that peak metal inputs occur in the most recent period.
426 Enrichment factors (Fig. 3) of trace metals also show little change, remaining around 1.0 throughout
427 the dated periods for both cores indicating no clear inputs from anthropogenic sources. Again, the
428 only exception is Hg, which increases slightly from 1.0 in the 1970s to a maximum of 1.3 in the surface
429 sediments of the lake core and indicating recent minor enrichment (Cevik et al., 2009).

430

431 Concentrations of geochemical elements and trace metals (Mg, Al, Si, S, K, Ca, Ti, Mn, Fe, Br, Rb, Sr,
432 Ti, V, Cr, Ni, Cu, Zn, As, Pb, Hg) in the LETS2 sediment core along with LOI_{550} as an estimate of organic
433 matter were standardised (z-scores) prior to principal components analysis (PCA) to explore
434 relationships between these data. PCA axis 1 explained 38.5 % of the total variance, and described
435 high loadings of minerogenic elements and some trace metals (e.g. Al, Ti, Rb, Ni, Cu) against negative
436 loadings of Hg and Mn (Supp. Info. SI2). PCA axis 2 eigenvalues (28.7% variance) corresponded to the
437 difference in loadings between organic (LOI_{550} , Br) and minerogenic elements (Sr, K). Axis 3 explains
438 10.3% of the total variance and, along with axes 1 and 2, describes cumulatively 77% of the variation
439 in these element concentration data. Component weightings of the variables for the first 3 PC axes
440 are also shown in Supp. Info. SI2. Stratigraphically, the bi-plot is reflected in the co-increase of Hg with

441 organic matter and Mn against the relatively unchanging concentrations of other trace metals and
442 minerogenic-associated elements.

443 At Letšeng-la Letsie, spheroidal carbonaceous fly-ash particles (SCPs) are first recorded in the lake
444 sediment core in the late-1970s and in the wetland core in the late-1980s. Peak concentrations agree
445 well, reaching 900 gDM^{-1} in 2015 in the lake record and 500 gDM^{-1} in wetland surface sediments (Fig.
446 3). These SCP concentrations are similar to those found in recent levels of a peat core ($200 - 350 \text{ gDM}^{-1}$;
447 unpublished data) taken from Sani Top ($29^{\circ}35'8.12''\text{S}$; $29^{\circ}15'56.68''\text{E}$ - Fig. 1; Fitchett et al., 2017)
448 140 km to the northeast of Letšeng-la Letsie. For both cores, SCP flux records reinforce the
449 concentration patterns and trends (Supp. Info. SI2) in showing a steady increase in contamination
450 through to the present although recent fluxes in the wetland core ($326 \text{ cm}^{-2} \text{ yr}^{-1}$) are approximately
451 double those of the lake ($148 \text{ cm}^{-2} \text{ yr}^{-1}$).

452

453 The $\delta^{15}\text{N}$ data for both lake and wetland cores are also shown in Fig. 3. For the lake core, $\delta^{15}\text{N}$ values
454 range between 5.5 and 5.9 ‰ from basal sediments (1930s) through to the late-1970s, when they
455 increase gradually to maximum values of around 6.5 ‰ in most recent sediments. By contrast, $\delta^{15}\text{N}$
456 values in the wetland core are higher, remaining between 7.5 and 8.1 ‰ from the base of the core
457 (1890s) through to the early-2000s where they decline slightly, reaching a minimum of 6.7 - 7.0 ‰ in
458 surface sediments.

459

460 *3.3. Organic pollutants*

461

462 Concentrations of total PAHs, total 16 United States Environmental Protection Agency (USEPA) PAHs,
463 total high molecular weight (HMW; 4-6 rings) and total low molecular weight (LMW; 2-3 rings) PAHs
464 in LETS2 all follow very similar historical trends (Fig. 3c). Concentrations are low, with total PAH
465 concentrations not exceeding 250 ng g^{-1} . Concentrations for all these groups of compounds remain
466 largely unchanged through the core except for two minima around 1970 and the early-2000s, and a

467 peak in the late-1990s. The agreement between these different groups suggests similar sources for
468 both HMW and LMW compounds. However, concentrations for some individual HMW PAHs e.g.
469 benzo[*a*]pyrene, indeno[1,2,3-*c,d*]pyrene and benzo[*g,h,i*]perylene only exceed the limit of detection
470 after the mid-1970s, with concentrations increasing from this time through to surface sediments. High
471 molecular weight PAHs are preferentially industrially-derived (Tobiszewski and Namieśnik, 2012;
472 Thang et al., 2019) suggesting increasing inputs from these sources over the last 40 years. This is
473 confirmed by the flux data for these compounds which reaches a peak of 4.6 ng cm⁻² yr⁻¹ in surface
474 sediments. Concentration ratios of PAH compounds may also be used to provide information on
475 sources. The ratio between non-alkylated phenanthrene and total (alkylated and non-alkylated)
476 phenanthrene ($C_0\text{-P}/(C_0 + C_1)\text{-P}$) and the ratio between fluoranthene and the sum of fluoranthene and
477 pyrene ($\text{Fl}/\text{Fl} + \text{Py}$) can both provide an indication as to the strength of the role played by various
478 combustion sources. For both ratios, values > 0.5 indicate combustion of biomass and / or coal (Yunker
479 et al., 2002) while those 0.4 - 0.5 indicate combustion of petroleum. Ratios exceed 0.4, and mostly 0.5,
480 throughout the core (Suppl. Info. SI3) indicating combustion as the only detected source of PAHs.

481

482 Concentrations of total PCBs and the total of the 7 'indicator' congeners (PCBs 28, 52, 101, 118, 138,
483 153, 180; hereafter PCB₇) are both low and only increase to above detection limits in the 1970s (Fig.
484 3c). As indicators of anthropogenic industrial contamination, independent of PAHs, these data provide
485 additional evidence for deposition from industrial sources at Letšeng-la Letsie. Concentrations of total
486 PCBs and PCB₇ increase through to the early-2000s but then show neither significant increase nor
487 decrease in the last two decades. Of the polybrominated flame retardants, only BDE47 is detected in
488 Letšeng-la Letsie, but concentrations are close to analytical detection limits. A BDE47 'peak'
489 concentration of 0.05 ng g⁻¹ is observed in surface sediments. Hexachlorobenzene (HCB) is first
490 detected in sediments dated to the 1950s and highest concentrations and fluxes occur in surface
491 sediments (0.21 ng g⁻¹; 0.03 ng cm⁻² yr⁻¹).

492

493 **4. Discussion**

494

495 *4.1. The sediment record at Letšeng-la Letsie*

496

497 There has been some uncertainty regarding the presence of a permanent waterbody at Letšeng-la
498 Letsie prior to the 1968 impoundment. While a first edition map from 1973, based on South African
499 aerial surveys in 1964 shows no lake in the area, the earliest 1:50,000 topographic maps (Basutoland
500 Sheet 3028AC, surveyed in 1950 and first published by the Directorate of Colonial Surveys, 1952) show
501 the presence of a pond or tarn in part of the present lake. Furthermore, a British Army survey from
502 1904 describes a *“flat open valley about 3.5 miles long by 1.5 miles broad, surrounded by mountains.*
503 *There are a number of pools surrounded by boggy land; branch valleys contribute small streams which*
504 *combine to form the northern source of the Quthing River”* (Dobson, in South Africa 1910) (see Supp.
505 Info. SI4 for comparison).

506

507 Neither the wetland nor the lake sediment core indicate any evidence for a hiatus in accumulation
508 before or after the construction of the impoundment in 1968 and the radiometric chronologies for
509 both cores continue through this period. This is to be expected for the wetland core as the sampling
510 location is above the current water level and is therefore likely to have been so throughout the period.
511 Accumulation rates continue a smooth increasing trend before, during and after dam construction and
512 geochemical trends also exhibit no changes (Supp. Info. SI5). For the lake core (LETS2), sediment
513 accumulation rates are more variable prior to 1968 (Supp. Info. SI1) suggesting a more stable
514 accumulating environment after the impoundment, while there is also a marked, although brief,
515 decline in Fe/Mn ratio at the time of its construction (Supp. Info. SI5). Fe/Mn declines in reducing
516 conditions and it maybe that this temporary dip is due to extended inundation of the catchment when
517 the dam was built. However, the Fe/Mn ratio reverts to pre-dam levels after a single datapoint so
518 should be interpreted with caution (Supp. Info. SI5). A small dip in LOI₅₅₀, and a small peak in K, Mn

519 and Ca also occur, briefly, at this time. It is unclear how significant these changes are as greater, longer-
520 term and monotonic changes in Fe/Mn occur from the 1990s which, since the mid-2000s seems to be
521 driven by an increase in Mn concentration. It seems unlikely that this more extensive recent trend
522 represents a further increase in catchment inundation as the dam has not been raised and the existing
523 structure has been reported to be in a state of poor repair since at least the late-1980s (Skoftefeld,
524 1988). Furthermore, it is unlikely that this shallow, wind-stressed lake experiences extended periods
525 of anoxia even at its deepest point where the sediment core was taken. Organic matter content (as
526 LOI₅₅₀) also increases over this same period, but this is probably due to less diagenesis of organic
527 material in recent sediments. In summary, we conclude it most likely that a shallow lake or wetland
528 has been present at Letšeng-la Letsie over at least the last 85 years with continuous accumulation
529 both prior to, and following, the construction of the dam.

530

531 *4.2. Historical records of contamination*

532

533 The lake and wetland cores show continuous sedimentation over 85 ± 16 and 119 ± 23 years
534 respectively and thereby cover a period of increasing emissions within South Africa since the middle
535 of the 20th century. This includes the combustion of coal for the generation of electricity and other
536 industrial processes (Masekoameng et al., 2010). As Lesotho, by comparison, has only very minor
537 emissions and no industrial fossil-fuel combustion, we have hypothesised that any contamination
538 observed at Letšeng-la Letsie would likely be derived from the long-range transport of emissions from
539 South African sources.

540

541 Although the scale of contamination at Letšeng-la Letsie is very low, a number of measured lake
542 sediment and wetland parameters indicate an increase in contamination from the 1970s. While most
543 trace metals show no indication of anthropogenic enhancement, Hg concentrations and fluxes
544 increase from the 1970s, while the Hg EF also increases above 1.0 at this time albeit only reaching a

545 level of minor enrichment in surface sediments. Gaseous mercury emitted from industrial sources may
546 have an atmospheric lifetime of around one year and so Hg deposited at Letšeng-la Letsie is potentially
547 derived from anywhere globally (Yang et al., 2010a). However, some studies (e.g. Yang et al., 2016;
548 Panichev et al., 2019) indicate elevated Hg deposition at sites more proximal to emission sources. As
549 coal combustion is a major industrial Hg source, the coal-fired power plants in South Africa, which
550 contribute up to 78% of total Hg emissions in the country (27 - 39 tonnes yr⁻¹; Masekoameng et al.,
551 2010) are likely to be making a significant contribution to Hg deposition, especially given the timing of
552 the deposition increase. SCPs, also produced from the industrial combustion of coal, similarly show a
553 first presence and then a gradual increase from the 1970s in the lake sediment core and 1980s in the
554 wetland core. As these particulates have no natural sources, any presence may be considered an
555 enrichment, and the SCP concentration and flux profiles show good temporal agreement increasing
556 through to the present. SCPs are not able to travel globally prior to deposition, as Hg does, but may
557 travel thousands of kilometres under favourable meteorological conditions (Rose et al., 2012) and well
558 within the range required for transport from coal combustion sources in the Highveld (Fig. 1). The
559 agreement between Hg and SCP records provides a strong indication of South African industrial coal
560 combustion being the source of these contaminants.

561

562 The records of industrial organic contaminants and $\delta^{15}\text{N}$ also support this conclusion. The PAH ratio
563 data (Supp. Info. SI3) indicate that their source has been from combustion over the last century.
564 However, while historical patterns of total PAH concentrations vary little through time, those of
565 individual HMW compounds only exceed their analytical detection limits from the 1970s suggesting
566 an increase in emissions from industrial combustion sources. Similarly, the concentrations of total
567 PCBs and PCB₇ (Fig. 3) have only become elevated above detection limit since the 1970s, providing
568 independent evidence for the long-range transport and deposition of industrial emissions. Declining
569 trends in $\delta^{15}\text{N}$ recorded in lake sediment and ice cores (Holtgrieve et al., 2011) are considered
570 indicative of nitrogen from anthropogenic sources, including industrial fossil-fuel combustion, and so

571 the $\delta^{15}\text{N}$ decline over the last two decades in the Letšeng-la Letsie wetland core may reflect an increase
572 in atmospherically deposited nitrogen inputs from these sources (Heaton, 1987). The contrasting
573 increase in $\delta^{15}\text{N}$ in the lake sediment core since the 1970s clearly does not support this. However, it
574 may result from the expansion of the waterbody following the impoundment in 1968 and an
575 enrichment from animal-derived nitrogen inputs from newly flooded areas entering the sediment
576 record from this time.

577

578 There is clear evidence in the natural archives of Letšeng-la Letsie for an increase in the products of
579 high temperature coal combustion since the 1970s and this reflects the development of the power
580 generation and other heavy industries in South Africa. The 1970s saw a marked increase in electricity
581 generation in South Africa with the commissioning of a number of new, and larger, coal-fuelled power
582 stations mostly in the Highveld / Mpumalanga region. The amount of coal burnt in coal-fired power
583 stations more than doubled from 22 Mt in 1970 to 47 Mt in 1980, reaching a peak of 125 Mt in 2007.
584 This resulted in concomitant increases in the emissions of SO_2 and Hg. Sulphur dioxide emissions
585 increased from 250 ktonnes in the late-1960s to a peak of almost 2000 ktonnes in 2013, while Hg
586 emissions increased from 4.5 tonnes in 1970 to a peak of over 22 tonnes yr^{-1} between 2007 and 2012
587 (Fig. 4). Importantly, while coal combustion for electricity generation is the major source of Hg
588 emissions, other sources such as coal gasification, cement production and domestic coal burning also
589 contribute and increased through this period (Masekoameng et al., 2010). However, as the scale of
590 contamination is low, further evidence is required to track these contaminants to likely source areas.

591

592 *4.3. Tracing the sources of contamination*

593

594 The deposition of contaminants in the Maloti Mountains of southern Lesotho from suspected
595 industrial sources in the Highveld of South Africa requires long-range transport. Back-trajectory
596 analysis permits the reconstruction of air-mass pathways to a location, and is widely used to

597 determine potential sources in contaminant studies. Air originating over the industrialised Highveld
598 and the Lephalale region is advected over Letšeng-la Letsie, but only infrequently (no more than 7%
599 of the time; Fig. 5) which may explain the low level of contamination experienced at the site. This flow
600 occurs in association with continental anticyclones and the slow-moving air-masses typically take 2-3
601 days to move from the Highveld to southern Lesotho. By contrast, air masses transported from the
602 south-west and west (occurring 56% of the time), from the east (10%), and from Botswana in the north
603 (26%) are relatively free of contamination from coal combustion. These flow patterns have been
604 broadly similar at least since 1950 (Supp. Info. SI6 for more detailed historical breakdown). Aerosols
605 from industries and biomass burning have previously been measured at Ben Macdhui in Eastern Cape,
606 South Africa, about 45 km to the south-south-west of Letšeng-la Letsie (Piketh et al., 1999; 2002).
607 Trajectory analysis at Ben Macdhui show that peak concentrations of industrial sulphate aerosols
608 occur in association with recirculating, anticyclonic flow which transports pollutants from the
609 Mpumalanga Highveld (Zunckel et al., 1999; Piketh et al., 2002).

610 Other industrial contaminants, such as PCBs, which show some coincidence with the historical
611 patterns of PAHs, may also be derived from industries in the Highveld region and an assessment of
612 the PCB congeners helps corroborate our hypothesis that these are derived from long-distance
613 transport. PCBs congener distribution was dominated by low-chlorinated forms (fewer than four
614 chlorine atoms), which are more volatile and mainly related to atmospheric transport. Furthermore,
615 the ratio $R = ([PCB28] + [PCB52]) / ([PCB28] + [PCB52] + [PCB138] + [PCB153] + [PCB180])$ has been used
616 as an indicator of contaminant transport, where an R value close to 1.0 is considered to be related to
617 atmospheric deposition and long-range transport (Naffrechoux et al., 2015). Here, all calculated R-
618 values were 1.0, as the high-chlorinated forms included in this ratio were not detected. Similarly, the
619 flame retardant BDE47 recorded at the site, is most likely to be derived from urban or commercial
620 sources and so may be transported to Letšeng-la Letsie along the same atmospheric pathways. In
621 southern Africa, global warming is associated with a poleward shift of easterly and westerly wave
622 circulation systems (Sousa et al., 2018), an intensification of subtropical anticyclones, and heightened

623 variability associated with El Niño-Southern Oscillation (ENSO), the Indian Ocean Dipole (IOD) and the
624 Southern Annular Mode (SAM) (Fauchereau et al., 2003; Reason and Roault, 2005; Gaughan et al.,
625 2016). Over time, Letšeng-la Letsie is expected to be less frequently influenced by westerly flow and
626 more frequently influenced by anticyclonic flow, which may result in a marginal increase in transport
627 of polluted air from the industrialised regions to the north.

628

629 Hexachlorobenzene (HCB), was widely used as a fungicide but banned under the Stockholm
630 Convention on Persistent Organic Pollutants in 2001. It is a byproduct of the manufacture of certain
631 industrial chemicals and exists as an impurity in several pesticide formulations (UNEP, 2002). As
632 production and use of HCB have been regulated over the past decades an important source of this
633 contaminant is its unintentional release as a byproduct in many chlorination processes in chemical
634 and metals manufacturing. It may also be re-emitted from soils previously contaminated by
635 agricultural use as well as combustion processes including coal burning (Bailey, 2001; Gong et al.,
636 2017). Hexachlorobenzene is found consistently through the lake sediment core since the mid-1970s
637 (Fig. 3). A review of organochlorine compounds in South African freshwaters (Ansara-Ross et al., 2012)
638 demonstrated very few data for HCB except some sediment and soil concentrations in industrial and
639 agricultural areas (Quinn et al., 2009). Its presence has also been detected in the eggs of aquatic, and
640 aquatic-feeding, birds (Bouwman et al., 2008) and some freshwater fish species (Wepener et al., 2012;
641 Gerber et al., 2016) and Fatoki and Awofolu (2004) reported high concentrations (up to 150 ng L⁻¹) in
642 the waters of the Buffalo River in South Africa's Eastern Cape to the south-east of Lesotho.
643 Hexachlorobenzene has been banned in South Africa since 1983 (Bouwman, 2004) and its use is
644 considered limited (Gerber et al., 2016). Potential sources for HCB in Letšeng-la Letsie are, therefore,
645 difficult to identify, although Ansara-Ross et al. (2012) cite accidental spillages, poor management of
646 stocks and poor disposal as possible sources to the environment, in addition to its use for agricultural
647 purposes. Fatoki and Awofolu (2004) suggest that illegal use of pesticides may also be a source. One
648 agricultural application is as a scabicide treatment for sheep, which has been reported elsewhere in

649 Africa (e.g. Jemaa et al., 1986). Over 60% of households surveyed in the area around Letšeng-la Letsie
650 own grazing animals, including cattle sheep and goats, with sheep being the dominant form of
651 livestock (Lannas and Turpie, 2009). However, while nearly 90% of those surveyed grazed their stock
652 in the wetland area, we have found no evidence that HCB was used as a scabicide there. Another
653 possible source for HCB is as a by-product of coal combustion. Hexachlorobenzene is more volatile
654 than many POPs and is stable in the atmosphere. Consequently, it has a high potential for long-range
655 transport (Jaward et al., 2004; Shen et al., 2005; Pozo et al., 2017) and therefore HCB at Letšeng-la
656 Letsie could be derived from Highveld coal emissions. Further research is certainly warranted to
657 investigate organochlorine compound patterns in the region.

658

659 *4.4. Scale of contamination*

660

661 Contamination measured in Letšeng-la Letsie archives is low with many POPs at or below the analytical
662 limit of detection. All trace metal concentrations, except those of Ni and Cu, are well below their
663 consensus threshold effect concentration (TEC), the concentration at which any detrimental biological
664 effect would be expected to occur only rarely (MacDonald et al., 2000; Buchman, 2008).
665 Concentrations of both Ni and Cu are approximately double their TECs (22.7 and 31.6 $\mu\text{g g}^{-1}$
666 respectively) throughout the lake sediment core and Ni concentrations approach the Probable Effects
667 Concentration (PEC) (48.6 $\mu\text{g g}^{-1}$), the level at which detrimental biological effects would be expected
668 due to that contaminant alone. However, both Cu and Ni have EFs of around 1.0 throughout the core
669 indicating that these elevated concentrations are most likely natural, as reported for other studies in
670 South Africa (Dalton et al., 2018). By contrast, and for comparison, the peak concentration of HCB in
671 the core is 0.21 ng g^{-1} , two orders of magnitude below the US National Oceanic and Atmospheric
672 Administration (NOAA) lowest effect level (LEL; no TEC reported for HCB) of 20 ng g^{-1} for freshwater
673 sediments. This is defined as the level of sediment contamination that can be tolerated by the majority
674 of benthic organisms (Buchman, 2008).

675

676 Comparing coal combustion-related contaminants in Letšeng-la Letsie sediments with those from
677 other remote lakes, the peak SCP flux of $148 \text{ cm}^{-2} \text{ yr}^{-1}$ is up to an order of magnitude higher than
678 equivalent fluxes in Greenland (Bindler et al., 2001) and in Antarctic or sub-Antarctic lake sediments
679 (Rose et al., 2012). They are at a similar level to mountain lakes in Europe (e.g. Rose et al., 1999)
680 (typically $100 - 180 \text{ cm}^{-2} \text{ yr}^{-1}$) and the Rocky Mountains of North America (Landers et al., 2008) ($50 -$
681 $100 \text{ cm}^{-2} \text{ yr}^{-1}$). While equivalent data in Africa are sparse, the peak SCP sediment flux for Lake Bisoke,
682 a volcanic crater lake in the Albertine Rift in Rwanda was similar ($179 \text{ cm}^{-2} \text{ yr}^{-1}$; McGlynn et al., 2019),
683 but higher than peak fluxes in an Ethiopian mountain lake (approximately $50 \text{ cm}^{-2} \text{ yr}^{-1}$; Bittner et al.,
684 in press).

685

686 The Hg peak concentration and EF in the Letšeng-la Letsie lake sediment core are 51 ng g^{-1} and 1.3
687 respectively. This concentration is equivalent to those reported from remote lakes on the Tibetan
688 Plateau ($30 - 50 \text{ ng g}^{-1}$; Yang et al., 2010a) but lower than Hg concentrations in sediments taken from
689 lakes in the Rwenzori Mountains in Uganda ($100 - 250 \text{ ng g}^{-1}$) (Yang et al., 2010b) and the US Rocky
690 Mountains ($100 - 550 \text{ ng g}^{-1}$) as well as lakes from central Alaska ($110 - 150 \text{ ng g}^{-1}$) (Landers et al.,
691 2008). The comparison for PAH contamination is similar to that of Hg, with peak concentrations of
692 total PAH and total 16 USEPA PAHs (244 ng g^{-1} and 126 ng g^{-1} respectively) similar to those for Qinghai
693 Lake on the northern Tibetan Plateau ($15-341$ and $11 - 279 \text{ ng g}^{-1}$ respectively) (Wang et al., 2010) and
694 in the same broad range as those reported for marine sediments in the vicinity of Antarctic islands
695 (Sutilli et al., 2019). These PAH concentrations are lower than those for lakes in the Rocky Mountains,
696 where total concentrations for the 16 USEPA PAHs in Lone Pine Lake and Mills Lake were reported as
697 $210 - 280 \text{ ng g}^{-1}$ (Usenko et al., 2007) and for European mountain lakes where peak total PAH
698 concentrations varied between 300 ng g^{-1} for Arresjoen on Svalbard and up to $25,000 \text{ ng g}^{-1}$ for remote
699 mountain lakes in the Tatra Mountains of Slovakia and Poland (Fernandez et al., 2002). Interestingly,
700 these latter concentrations are similar ($9700 - 61,000 \text{ ng g}^{-1}$) to those reported for soils taken from

701 the vicinity of coal-fired power stations in South Africa (Okedeyi, 2013) indicating elevated levels of
702 proximal PAH deposition close to sources.

703

704 It is widely accepted that Hg in the global atmosphere at the start of the 21st century had increased by
705 around 3-fold since pre-industrial times although more recent estimates place this increase at 450 %
706 (UNEP, 2019). Hence, Hg enrichments (ratios of modern to background but not necessarily calculated
707 as EFs) in remote lake sediments, not impacted by local or regional sources, are often seen to be
708 around these values. For example, in a review, Fitzgerald et al. (1998) reported enrichments of 2.7 (\pm
709 standard deviation of 0.9) for lakes in mid-continental North America, 2.3 (\pm 0.6) in northern Canada
710 and 2.0 - 2.6 in Scandinavia. Similarly, Yang et al. (2010a;b) show Hg enrichments of 2.5 - 2.9 for
711 Rwenzori Mountain lakes and 2.8 - 14.7 (the latter closer to sources) for Tibetan Plateau lakes. The EF
712 for Letšeng-la Letsie is considerably lower than these remote lakes and even selecting the highest and
713 lowest Hg fluxes only provides an enrichment ratio of below 2.0. Although low Hg values have been
714 reported for southern Africa in the past, leading to the suggestion of an African mercury “anomaly”
715 (e.g. Black et al., 2011), this has tended to be related to unexpectedly low Hg levels in biota which
716 would have been expected to biomagnify Hg through the aquatic food-web. However, given the
717 atmospheric life-time of Hg may conceivably be around one year and hence Hg deposited to remote
718 lake sediments could potentially be derived from any global source, an African anomaly may be less
719 relevant. Indeed, such an anomaly might be expected to result in lower Hg concentrations in
720 uncontaminated background sediments, which would then elevate Hg EFs in contemporary samples.
721 This is clearly not the case for Letšeng-la Letsie. It is currently unclear why Hg enrichment at the site
722 is lower than other global ‘background’ locations.

723

724 *4.5. Wider implications*

725

726 Estimates of anthropogenic Hg emissions to the atmosphere in South Africa increased in the first

727 decade of the 21st century from approximately 34 tonnes yr⁻¹ in 2000 to around 50 tonnes yr⁻¹ in 2006
728 (Masekoameng et al., 2010). Since then, coal combustion has levelled-off and is predicted to decline
729 slightly through to 2030. However, coal remains in South Africa's power generation plans at least
730 through to 2050 (South Africa Department of Energy, 2019) although there is currently no targeted
731 mercury abatement for the country. Coal combustion in the power generation industry remains the
732 largest Hg emission source by far in South Africa and so while predicted Hg emissions are also expected
733 to decline slightly to 2026 (Garnham and Langerman, 2016), emissions of this and other contaminants
734 from coal combustion such as PAHs, SCPs and other particulate matter are likely to continue into the
735 late-21st century being distributed and deposited throughout southern Africa. Currently, deposition of
736 these contaminants in Lesotho would appear to be limited, although comparable data within the
737 country, as for the rest of southern Africa, are remarkably sparse and so it is difficult to ascertain the
738 full scale and extent across the region.

739

740 Lesotho is an important water resource area for southern Africa and an estimated 780 million m³ of
741 water each year is exported to South Africa under the Lesotho Highlands Water Project (LHWP)
742 programme. Despite this, many rural communities still take drinking water directly from springs and
743 wells while urban water supplies experience frequent water scarcity. Maintaining and expanding
744 access to clean and secure water is one of the biggest challenges facing southern Africa in the 21st
745 century. This is likely to be exacerbated by the predicted impacts of climate change on the quantity
746 and timing of rainfall, and increases in temperature, making this area one of the world's most
747 climatically vulnerable regions. Determining the scale of contamination in these source areas is
748 therefore vital for both human and ecological health.

749

750 Furthermore, the health of aquatic ecosystems is recognised as an essential prerequisite for
751 sustainable use of water resources and Claassen (2010) states that: *"The biggest threat to a*
752 *sustainable water supply in South Africa is not a lack of storage but the contamination of available*

753 *water resources through pollution*". Further deterioration of water quality through increasing levels
754 of contamination could have a serious impact on ecosystem services provided by southern African
755 freshwaters, especially to rural populations. While the situation is already critical, the synergistic
756 effects of increased contaminant loading and climate-related stress are likely to lead to further
757 degradation in water quality and aquatic ecosystems. Such deterioration will also jeopardise the
758 suitability of natural waters for human use (e.g. contaminant intake through fish consumption)
759 (Wepener and Chapman, 2012) and adversely affect ecosystem services for rural communities.
760 Despite this, the scale, extent and rate of freshwater contamination are largely unknown and there is
761 an urgent need to address this. In the absence of monitoring, the analysis of natural archives such as
762 lake and wetland cores is one of the only ways by which to determine trends in the direction (i.e.
763 improvement or deterioration) and rates of change in freshwater contamination. The records
764 presented here which, to our knowledge, are the first high resolution multi-pollutant sediment records
765 for southern Africa therefore provide a valuable baseline, not just for temporal changes in deposition
766 to this background region, but also geographically for comparison to future studies in more
767 contaminated areas of southern Africa.

768

769 **Authorship contribution statement**

770 **Neil Rose:** Conceptualization; Data curation; Formal analysis; Funding acquisition; Investigation;
771 Methodology; Project administration; Visualization; Writing - original draft; Writing - review & editing.

772 **Alice Milner:** Conceptualization; Data curation; Funding acquisition; Investigation; Methodology;
773 Writing - review & editing. **Jennifer Fitchett:** Conceptualization; Funding acquisition; Investigation;

774 Project administration; Methodology; Project administration; Visualization; Writing - review & editing.

775 **Kristy Langerman:** Formal analysis; Methodology; Visualization; Writing - review & editing. **Handong**

776 **Yang:** Formal analysis; Validation; Writing - review & editing. **Simon Turner:** Data curation; Formal

777 analysis; Validation; Writing - review & editing. **Anne-Lise Jourdan:** Formal analysis; Writing - review &

778 editing. **James Shilland:** Formal analysis; Writing - review & editing. **César Martins:** Data curation;

779 Formal analysis; Methodology; Validation; Visualization; Writing - review & editing. **Amanda Câmara**
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781 Conceptualization; Funding acquisition; Investigation; Project administration; Resources; Writing -
782 review & editing.

783

784 **Declaration of Competing Interest**

785 The authors declare no competing financial interest.

786

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